

## Abstract

Rapid gravitational mass movements, such as landslides, rockfalls, or avalanches are repeatedly recognized during routine seismic monitoring at national earthquake observatories. Yet, utilizing the tools of seismology for fast detection and characterization of mass movements is still uncommon.

Seismology recently gained attention for the assessment of gravitational mass movements because it can potentially provide continuous realtime detection and approximate localization of events. However, the applicability on country scale needs to be tested and the seismological determination of precise location and of event parameters is challenging.

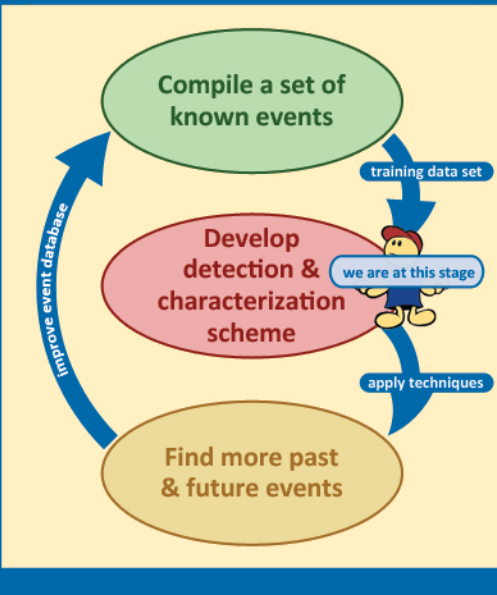
Here we present a set of past rockfall events in Austria and neighboring countries, which were well-recorded by several permanent and temporary seismic stations. We aim at identifying seismically observable parameters of rockfalls in comparison with additional geological and geographical data.

## Quick facts

dataset of **19 rockfalls**  
in/near **Austria** recorded  
between **2007 - 2016**

confirmed by **ZAMG**  
using **regional seismic network**  
(Zentralanstalt für Meteorologie und Geodynamik)

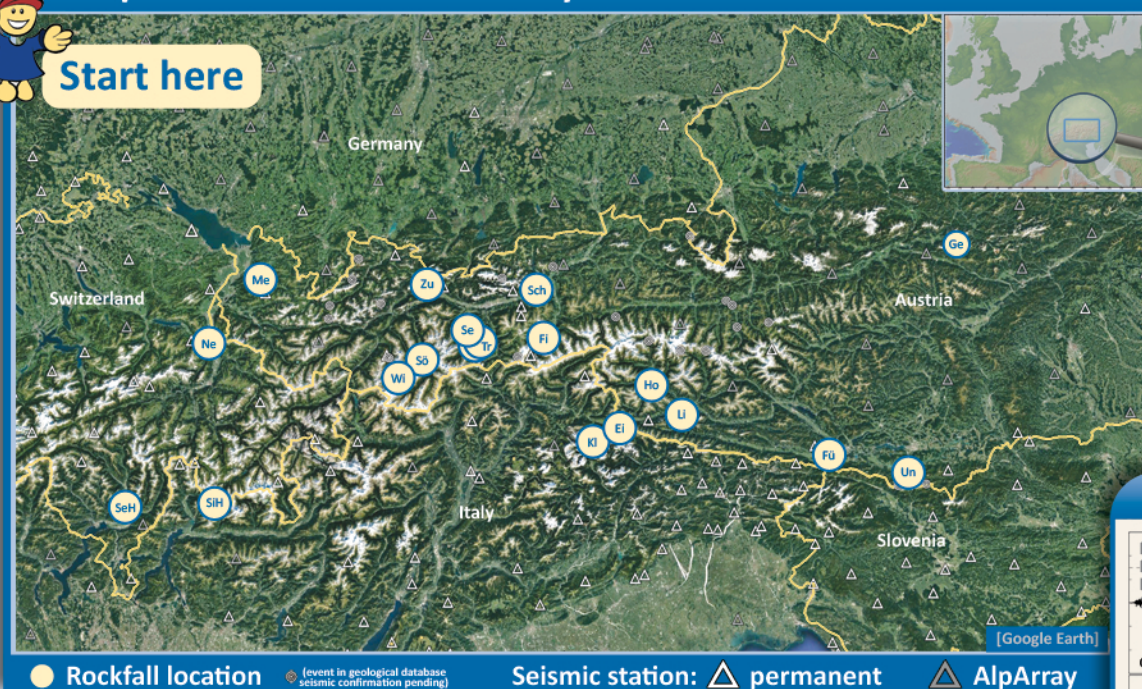
## Workflow



## Map of confirmed seismically detected rockfalls



Start here

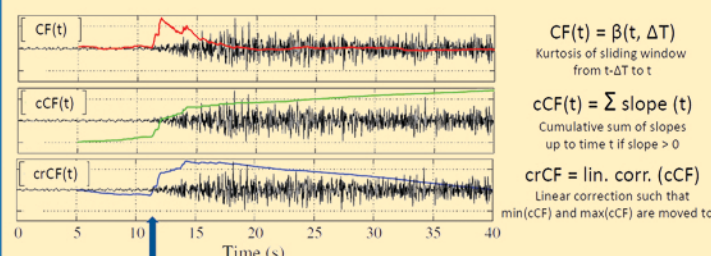


Detection with  
STA/LTA trigger

How to pick  
correct onset?

## Onset determination

For such emergent signals we determine signal onset from **Kurtosis change**  
(5 s sliding window for Kurtosis calculation around STA/LTA trigger time)



How well  
does it  
work?

## Comparison with true locations

The Kurtosis **onset picks** are treated as **Pg phases** and forwarded to a modified **HYPOCENTER** code (Lienert & Havskov) for proper location. Travel times are calculated using a **simple 1D velocity model**. No systematic outlier handling is applied at the current stage. **Deviation from true location:**

Seehorn:	0.7 km
Sölden:	4.3 km
Kleine Gaisl:	4.3 km
Finkenberg:	4.8 km
Mellental:	5.0 km
Unterbergen:	5.6 km
Nenzinger:	5.8 km
Zugspitze:	8.3 km
Einserkofel:	8.8 km
Silvretta:	9.0 km
Sellrain:	11 km
Gesäuse:	16 km
Fürnitz:	16 km
Trins:	26 km

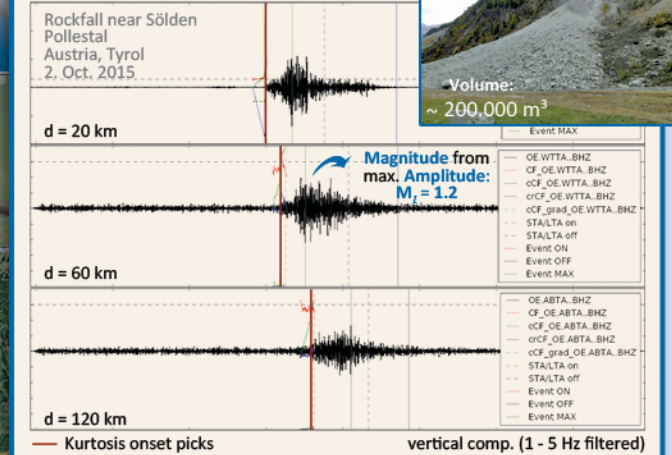
**Promising!**

Possible improvements:  
introduce error checks  
use local velocity model

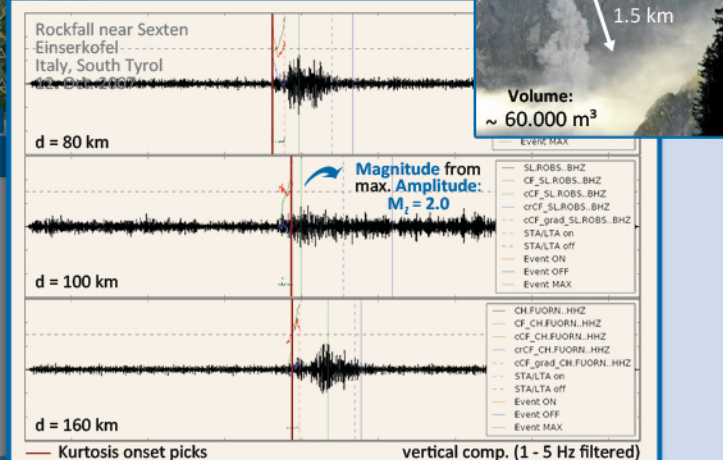
**Potential reason for deviation**  
(low SnR, STA/LTA activated on S wave)  
(low SnR, STA/LTA activated on S wave)  
(travel time too long for coincidence trigger)  
(low SnR, only 3 stations not well distributed)  
(bad SnR, picks wrong)  
(low SnR, station not well distributed)  
(low SnR, STA/LTA activated on S wave)

Kurtosis picks  
forwarded as  
Pg to HYPO-  
CENTER code

## Sölden, Austria



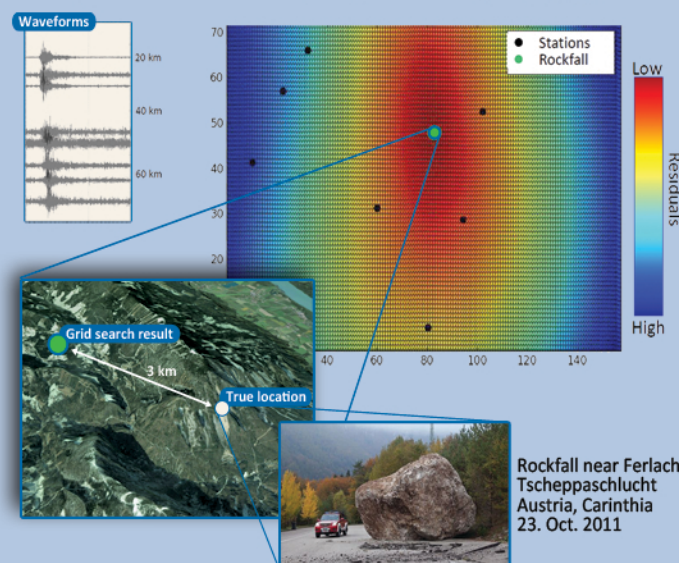
## Einserkofel, South Tyrol



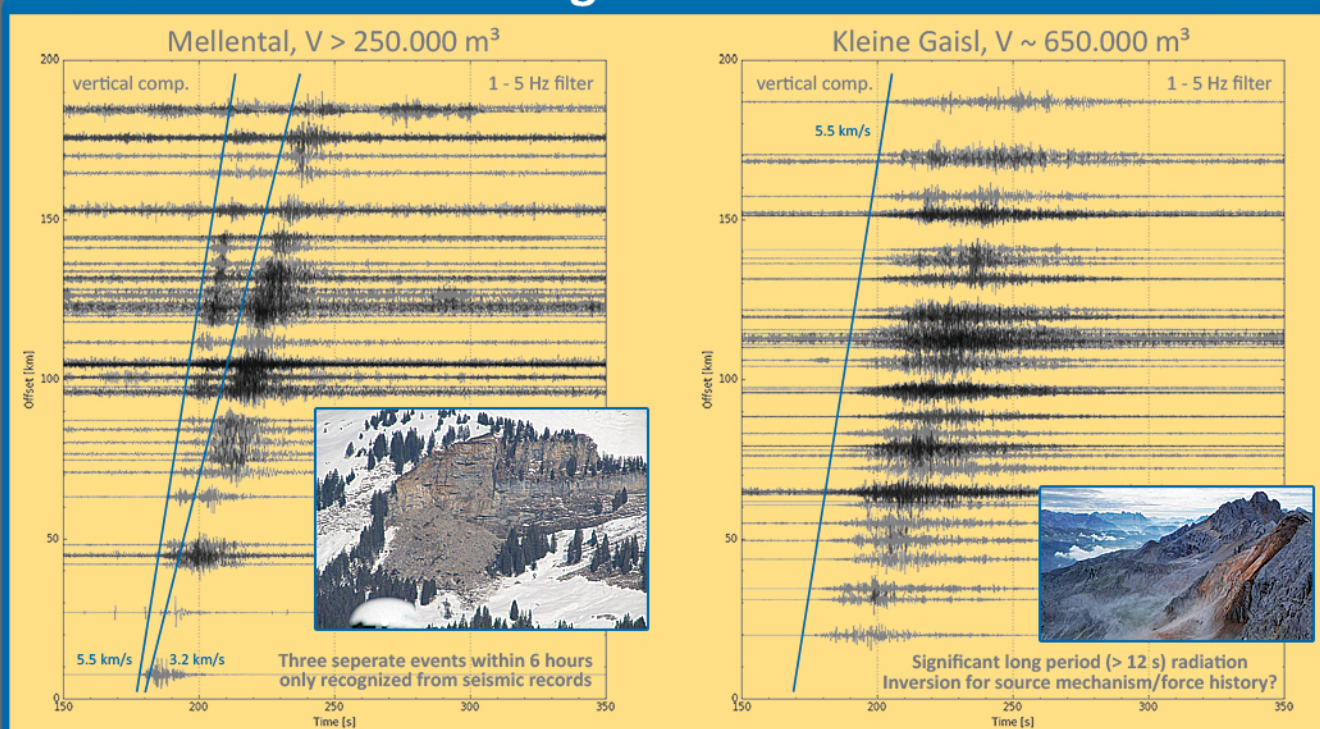
Is it precise enough for accurate location?

## Location

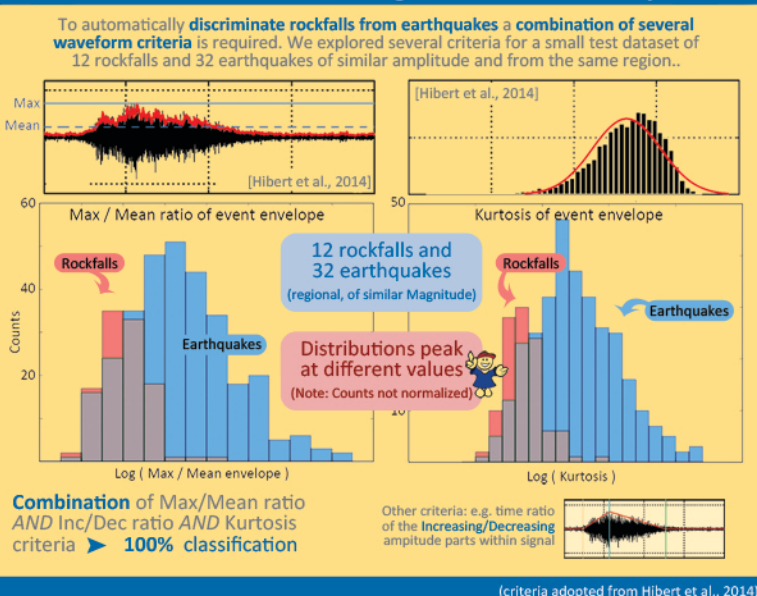
For **close stations** (< 100 km) preliminary location with a **grid search**  
(minimize travel time residuals, homogeneous velocity, no topography)



## Waveforms of two large events



## Discrimination from regional seismicity



## Conclusions (at this stage)

- Location of larger rockfalls on regional scale is possible  
Detection even works with comparably insensitive STA/LTA triggers
- Kurtosis based picking works even for low SnR  
(some error checking required to further improve accuracy)
- Rockfall/earthquake distinction by criteria combination  
Simple AND combination of 3 criteria sufficient for small test data set
- Seismic Magnitude ( $M_L$ ) does not relate to rock volume  
Which effect/mechanism is responsible for the maximum amplitude?
- Distant stations (> 100 km) can help to separate phases  
Open question: Which information is carried by the P- and S-waves?

## References

Hibert, C., Mangeney, A., Grandjean, G., Baillard, C., Rivet, D., Shapiro, N. M., Satriano, C., Maggi, A., Boissier, P., Ferrazzini, V. and Crawford, W. (2014), Automated identification, location, and volume estimation of rockfalls at Piton de la Fournaise volcano, JGR: Earth Surface, 119, 2014