AUTOMATIC SHEAR-WAVE SPLITTING ANALYSIS WITH SPLITWAVE

Gerrit Hein*, Eric Löberich, Götz Bokelmann

*gerrit.hein@univie.ac.at DEPARTMENT OF METEOROLOGY AND GEOPHYSICS, UNIVERSITY OF VIENNA

Abstract

The AlpArray Project is a joint effort to adress questions about Geodynamics in Central Europe. An important property in this context is seismic anisotropy. Shear-Wave splitting can be used to study the anisotropy of the eart. Previous studies ting axes in the Western Alps are nearly mountainchain parallel.

Many of the recent studies made use of the Split-Lab package by Wüstefeld et al. (2008).

Transverse minimisation (SC) and an Eigenvalue method (EV) from an open-source repository by Jack Walpole and compared these results with from the Chevrot-technique (2000). This poster will show preliminary SKS-results from often did this in a semi-automatic manner. Barruol the Swiss permanent network and demonstrate et al. (2011) for example could show, that fast split- the work flow for a single station. The goal is to establish a new automatic technique and apply it to the AlpArray network. The 600 stations recorded roughly the same amount of events as the permanent Swiss network did in 18 years.

Here, we deploy the Cross-Correlation (RC), a



Fig.1: Preliminary splitting parameters (time lag and fast direction) of SKS-splitting from eq. 2 at permanent Swiss stations.

Key facts:

- Swiss permanent network: 40 Stations
- Time window between: 2000-2018
- Number of Events $M_w > 7:240$
- Number of SKS-measurments: 3500
- Fixed time window around absolute maximum value $t_{win} = \pm 40$ s



Shear-Wave Splitting

An effect of anisotropy can be observed in Shear-Wave Splitting of SKS waves from teleseismic earthquakes. These waves travel as compressive waves through the outer liquid core and when passing through anisotropic layers divide into a fast and a slow pulse. The fast direction and the time lag are the the parameters which are usually sought when carrying out Shear-Wave splitting measurements. anisotropic layer



Fig.3.: Principle of S-wave splitting (Shearer, 2009).

There are different ways to obtain these parameters, e.g. the maximum cross-correlation (RC) method by Bowman and Ando (1987) and the transverse minimisation (SC) technique by Silver Chan (1991). These techniques allow determining fast direction and time delay for each station-event pair. The approach by Chevrot is somehow simpler. It assumes horizontal anisotropy and due to the similarity of components from the radially polarised SKS wave, it projects the transverse component (T) on the radial derivative (r) by a convolution with a splitting vector (s) (eq. 1). The splitting function describes the simple relation between the projected splitting intensities and the azimuthal event arrivals. This technique determines a single set of parameters per station, which is however relatively robust against low S/N ratios.

$$\mathsf{T}(t,\phi_R) = -\frac{1}{2}\mathsf{s} \circledast \mathsf{r} \tag{1}$$

where s varies as function of splitting time δ t and the azimuth of fast axis ϕ :

$$s(\phi_R) = \delta t \sin(\phi_R - \phi)$$

SKS splitting and splitting intensity



Fig.4: Example seismogram of an SKS-arrival.



For each station, the algorithm performs a grid search over a solution space. The fitted sinusoidal function in Figure 5 indicates the best fit for the splitting intensities and gives a robust estimate for δt and ϕ .





(2)

Fig.5: Splitting intensity plotted against the event azimuths together with best fit.



Fig.6: Misfit of delay-time and fast-axes between cross correlation method and the transverse minimisation method from Silver and Chan. The quality criteria were adapted after Wüstefeld et al. (2007).



Discussion and Outlook

- dependence of quality measurements.
- correction.

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Quality control

Preliminary method comparison

Relatively good agreement among methods. Manual results from Barruol can be reproduced.

Still needs thorough quality inspection. More investigations about time and frequency window

Future steps include a fully automatic Null detection, error surface stacking and a source-receiver